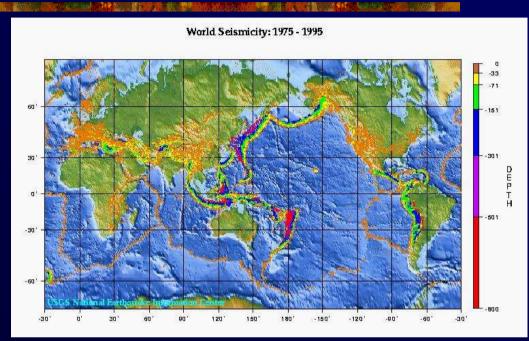
# Large-scale parallel simulations of earthquakes at high frequency: the SPECFEM3D project

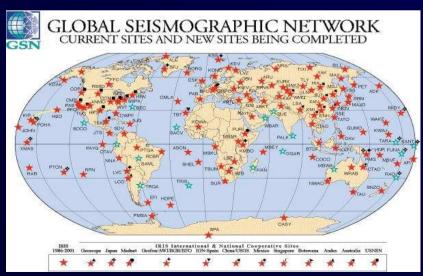
Dimitri Komatitsch, University of Pau, Institut universitaire de France and INRIA Magique3D, France
Jeroen Tromp et al., Caltech, USA
Jesús Labarta, Sergi Girona, BSC MareNostrum, Spain

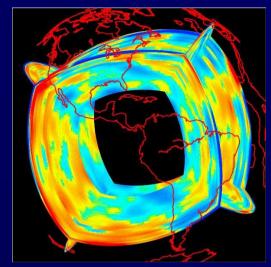
David Michéa, Nicolas Le Goff, Roland Martin, University of Pau, France

20 years of CERFACS, October 10, 2007

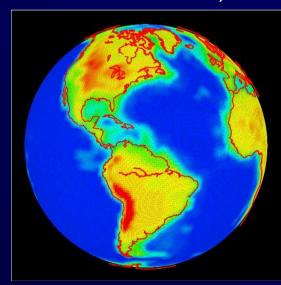
# Global 3D Earth







S20RTS mantle model (Ritsema et al. 1999)



Crust 5.2 (Bassin et al. 2000)

# Brief history of numerical methods

Seismic wave equation: tremendous increase of computational power

⇒ development of numerical methods for accurate calculation of synthetic seismograms in complex 3D geological models has been a continuous effort in last 30 years.

Finite-difference methods: Yee 1966, Chorin 1968, Alterman and Karal 1968, Madariaga 1976, Virieux 1986, Moczo et al, Olsen et al..., difficult for boundary conditions, surface waves, topography, full Earth

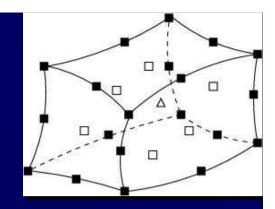
Boundary-element or boundary-integral methods (Kawase 1988, Sanchez-Sesma et al. 1991): homogeneous layers, expensive in 3D

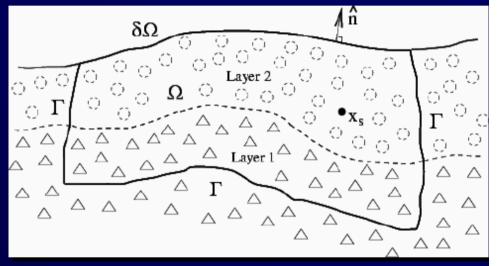
Spectral and pseudo-spectral methods (Carcione 1990): smooth media, difficult for boundary conditions, difficult on parallel computers

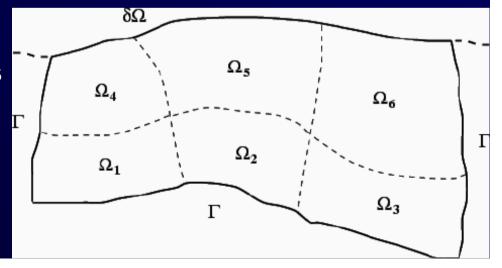
Classical finite-element methods (Lysmer and Drake 1972, Marfurt 1984, Bielak et al 1998): linear systems, large amount of numerical dispersion

# Spectral-Element Method

- Developed in Computational Fluid Dynamics (Patera 1984)
- Accuracy of a pseudospectral method, flexibility of a finite-element method
- Extended by Komatitsch and Tromp,
   Chaljub et al.
- Large curved "spectral" finiteelements with high-degree polynomial interpolation
- Mesh honors the main discontinuities (velocity, density) and topography
- Very efficient on parallel computers, no linear system to invert (diagonal mass matrix)







# Equations of Motion (solid)

Differential or strong form (e.g., finite differences):

$$\rho \, \partial_t^2 \mathbf{s} = \nabla \cdot \mathbf{T} + \mathbf{f}$$

We solve the integral or weak form:

$$\int \rho \mathbf{w} \cdot \partial_t^2 \mathbf{s} d^3 \mathbf{r} = -\int \nabla \mathbf{w} : \mathbf{T} d^3 \mathbf{r}$$

+ 
$$\mathbf{M} : \nabla \mathbf{w}(\mathbf{r}_s) S(t) - \int_{\mathsf{F}-\mathsf{S}} \mathbf{w} \cdot \mathbf{T} \cdot \hat{\mathbf{n}} \, d^2 \mathbf{r}$$

+ attenuation (memory variables) and ocean load

# Equations of Motion (Fluid)

Differential or strong form:

$$\rho \, \partial_t \mathbf{v} = -\nabla p$$

$$\partial_t p = -\kappa \nabla \cdot \mathbf{v}$$

We use a generalized velocity potential  $\chi$   $p = \partial_{\tau} x$ the integral or weak form is:

$$p = \partial_t x$$

$$\int \kappa^{-1} w \partial_t^2 x d^3 \mathbf{r} = -\int \rho^{-1} \nabla w \cdot \nabla x d^3 \mathbf{r}$$

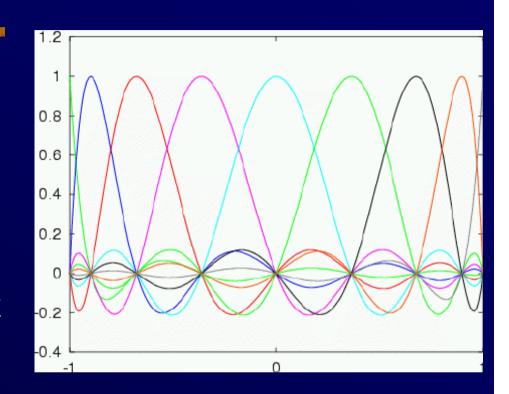
 $\Rightarrow$  3 times cheaper (scalar potential)

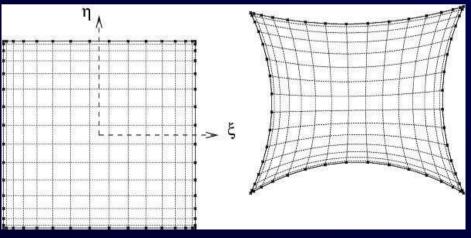
$$+\int_{\mathsf{F}-\mathsf{S}} w \hat{\mathbf{n}} \cdot \mathbf{v} \, \mathsf{d}^2 \mathbf{r}$$

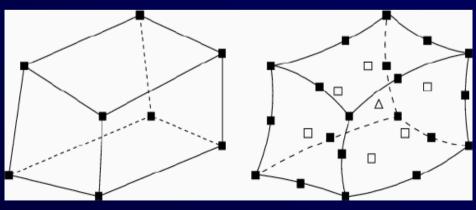
⇒ natural coupling with solid

## Finite Elements

- High-degree pseudospectral finite elements with Gauss-Lobatto-Legendre integration
- $\blacksquare$  N = 5 to 8 usually
- Exactly diagonal mass matrix
- No linear system to invert



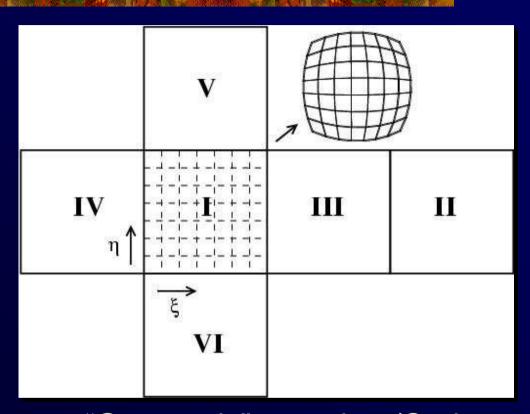


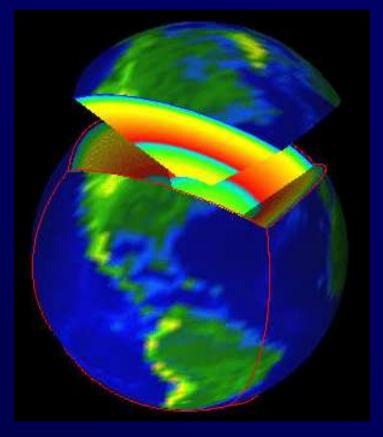


# The Challenge of the Global Earth

- A slow, thin, highly variable crust
- Sharp radial velocity and density discontinuities
- Fluid-solid boundaries (outer core of the Earth)
- Anisotropy
- Attenuation
- Ellipticity, topography and bathymetry
- Rotation
- Self-gravitation
- 3-D mantle and crust models (lateral variations)

# The Cubed Sphere

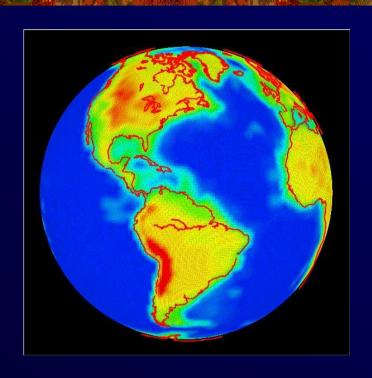


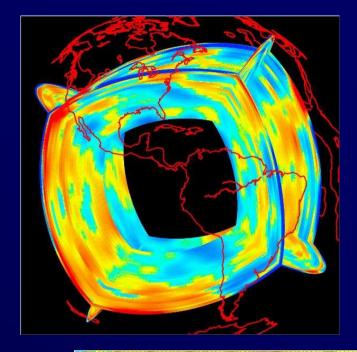


- "Gnomonic" mapping (Sadourny 1972)
- Ronchi et al. (1996), Chaljub (2000)
- Analytical mapping from six faces of cube to unit sphere

# Final Mesh

# Global 3-D Earth

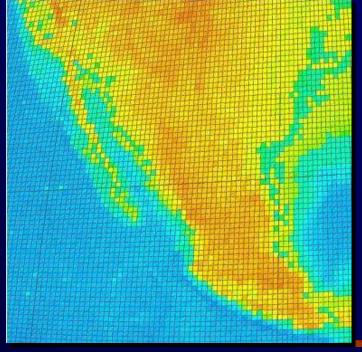




Crust 5.2 (Bassin et al. 2000) Mantle model S20RTS (Ritsema et al. 1999)

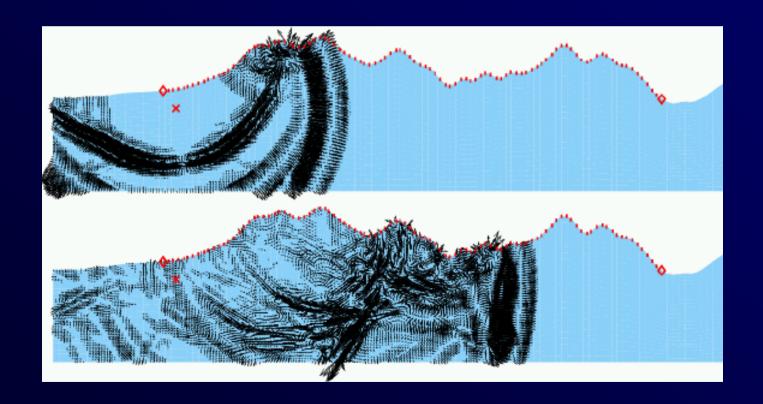
Ellipticity and topography

Small modification of the mesh, no problem

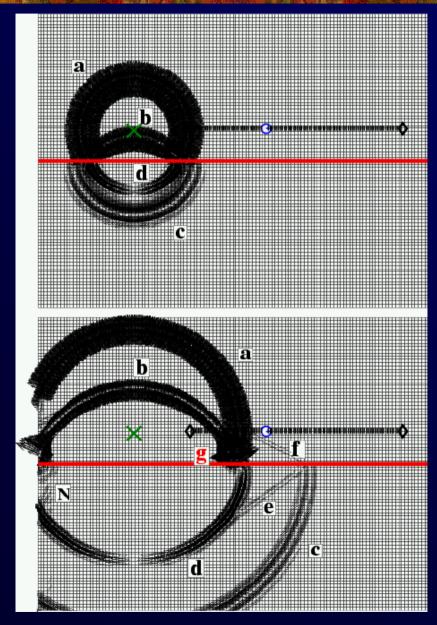


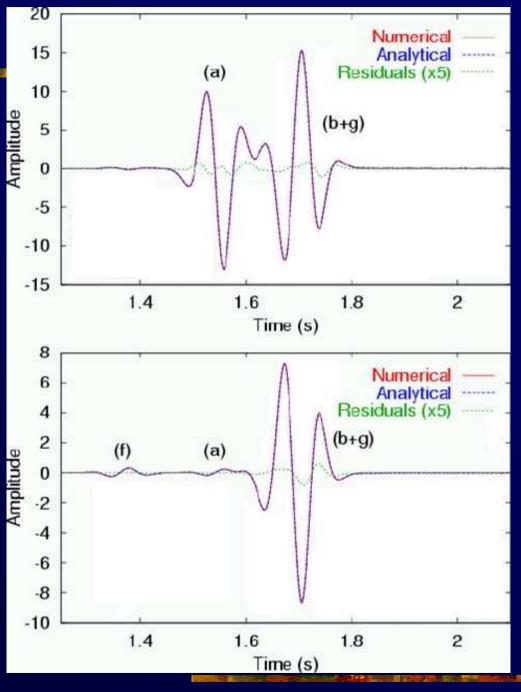
# Topography

- Use flexibility of mesh generation
- Accurate free-surface condition



# Fluid / solid



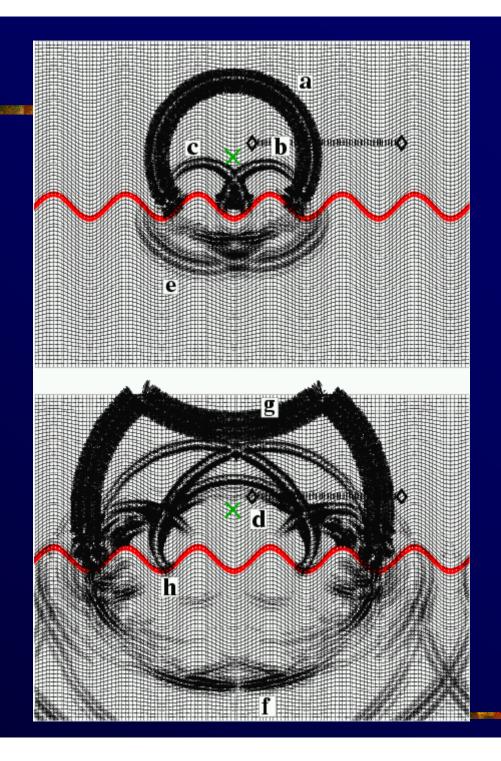


# Bathymetry

Use flexibility of mesh generation process

Triplications

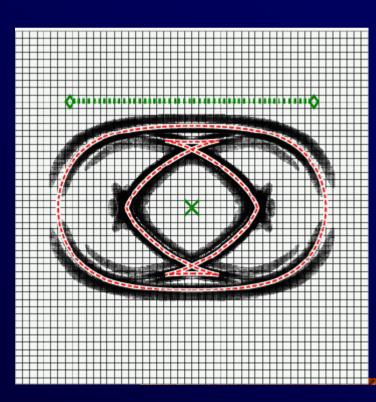
Stoneley



# Anisotropy

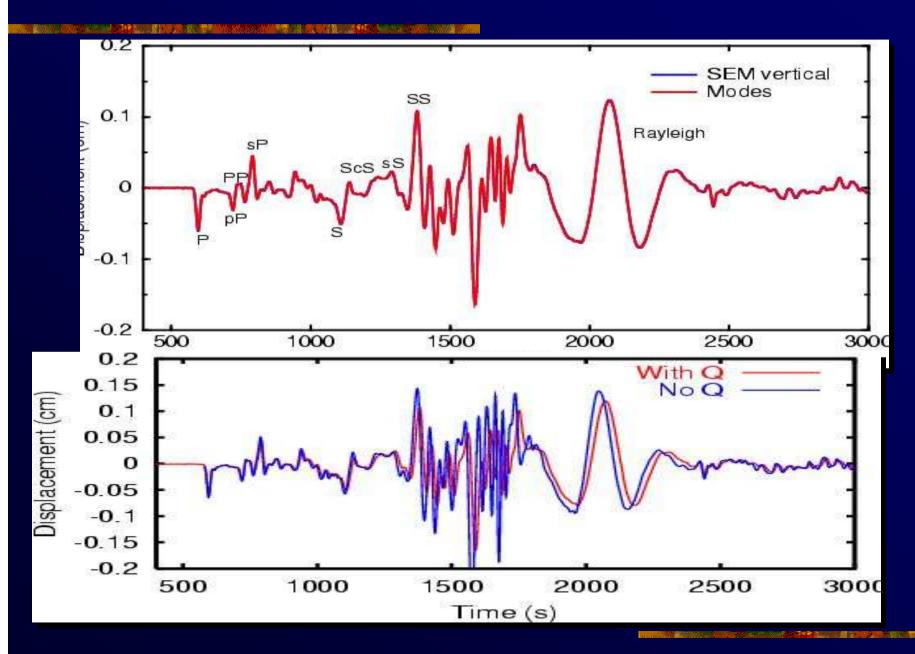
Cobalt

- Easy to implement up to 21 coefficients
- No interpolation necessary
- Tilted axes can be modeled

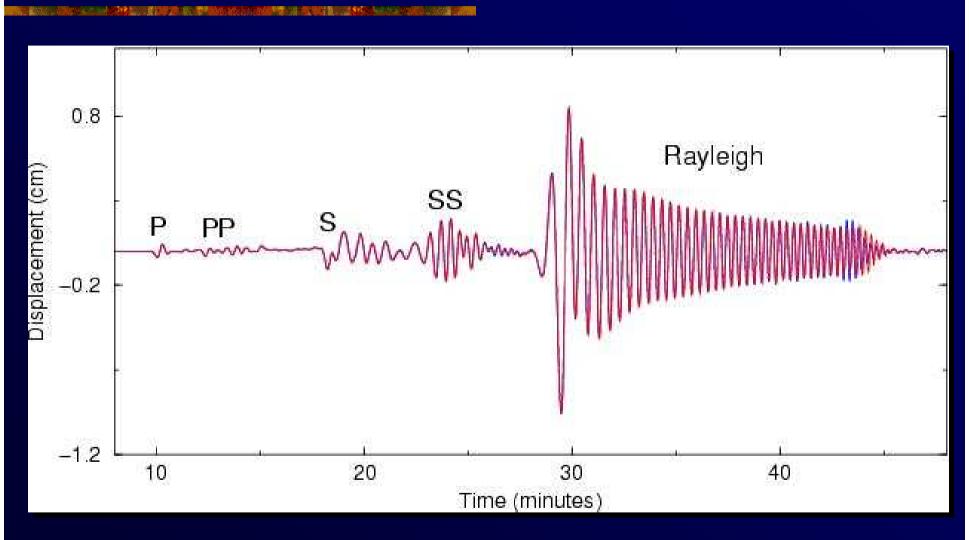


Zinc

# **Effect of Attenuation**



# Accurate surface waves

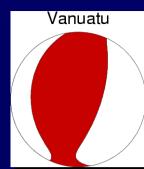


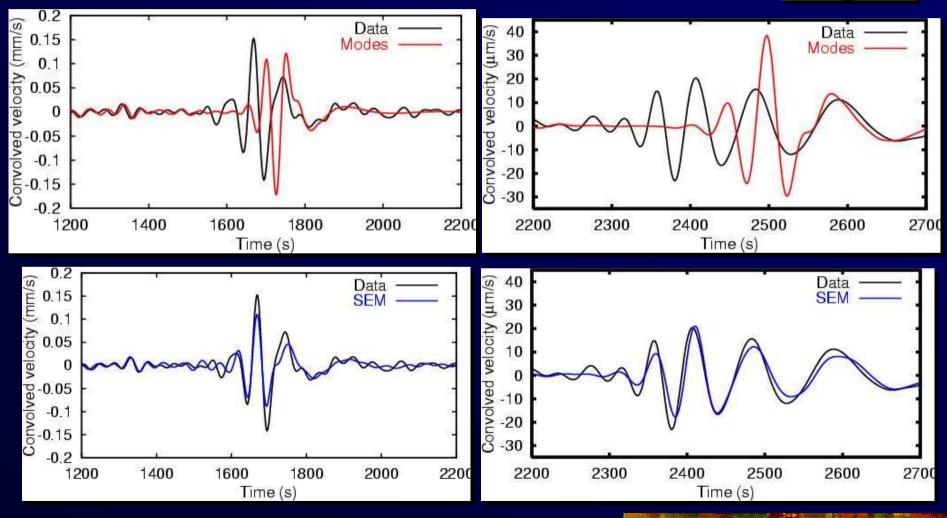
Excellent agreement with normal modes – Depth 15 km Anisotropy included

# Vanuatu

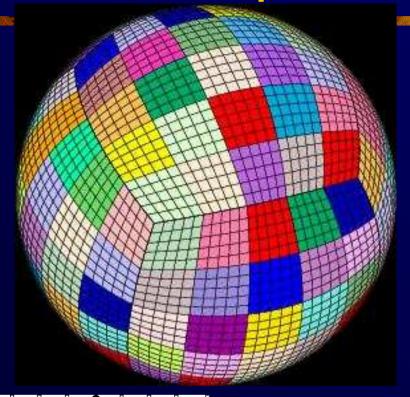
Depth 15 km

Composante verticale (onde de Rayleigh), trajet océanique, retard 85 s à Pasadena, meilleur fit au Japon

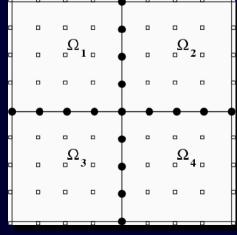




# Parallel Implementation

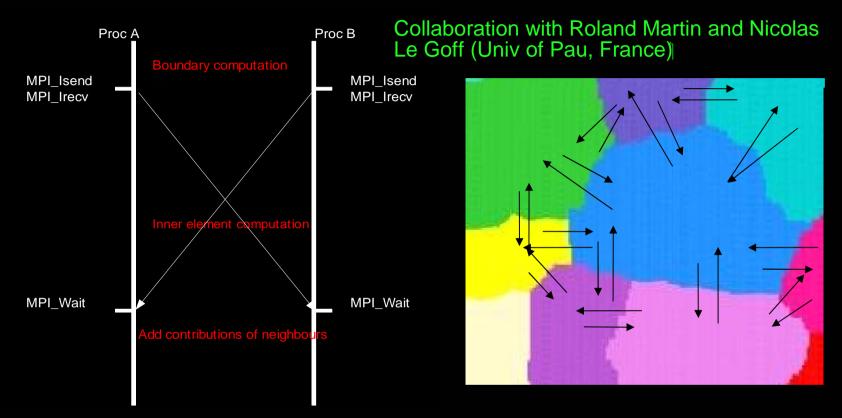






- Mesh decomposed into 150 slices
- One slice per processor MPI communications
- Mass matrix exactly diagonal no linear system
- Central cube based on Chaljub (2000)

#### Non-blocking MPI



Another way to optimize MPI code is to overlap communications with computations using non-blocking MPI. But, for our code, the overall cost of communications is very small (< 5%) compared to CPU time.

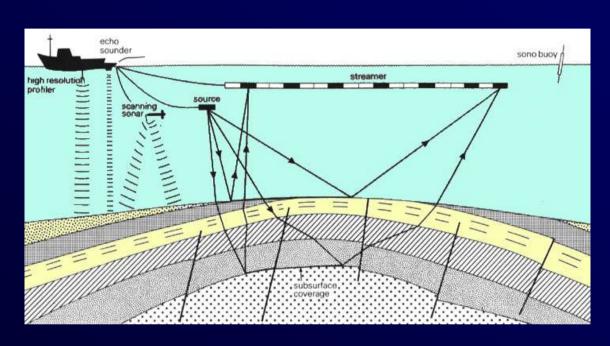
Also, looping on boundary elements contradicts Cuthill-McKee order and therefore causes cache misses.

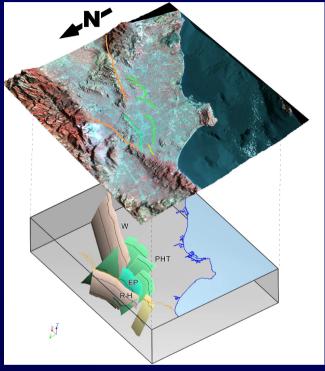
=> No need to use non blocking MPI because potential gain is comparable to overhead

=> Tested in 2D, and we did not gain anything significant

#### **Collaboration with the oil industry**





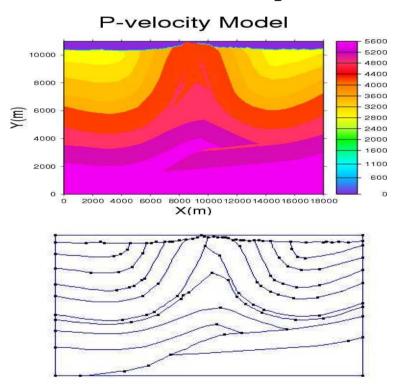


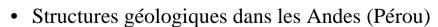
Dynamic geophysical technique of imaging subsurface geologic structures by generating sound waves at a source and recording the reflected components of this energy at receivers.

The Seismic Method is the *industry standard* for locating subsurface oil and gas accumulations.

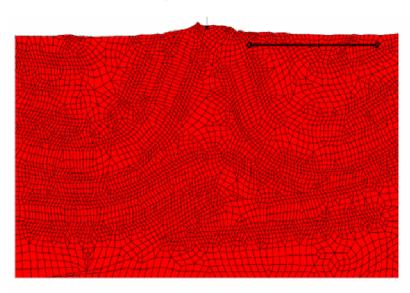
# Meshing an oil industry model

- Méthode d'éléments finis d'ordre élevé développée en **dynamique des fluides** (Patera 1984), en **sismique** 3Dpar Komatitsch et coll. (1998, 2002), Chaljub et coll. (2001).
- SPECFEM: Parallélisation MPI d'un Code F90 de 20000 lignes
  - → mailleur professionnel (GiD-UPC/CIMNE)





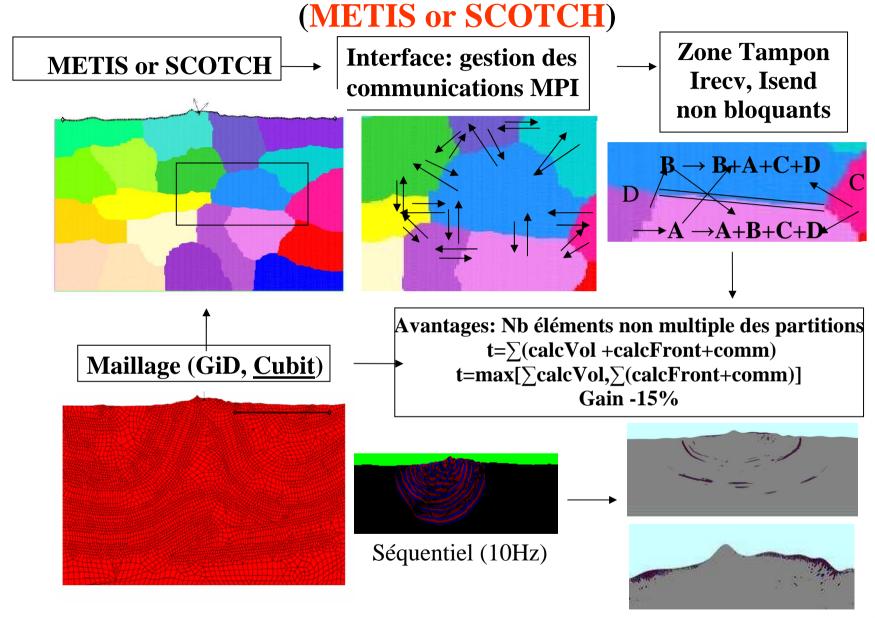
- Couche fine altérée en surface
- → Problème de dispersion en surface (Freq0 > 10 Hz).



- 5.3 millions de points à 10 Hz.
- Générateur GiD automatique de maillage (UPC/ CIMNE). 98% des angles 45° < θ < 135°. Pires angles: 9.5° and 172°

Sandrine Fauqueux Thèse INRIA/IFP (2003)

#### Partitionneur de domaine



#### Optimization of global addressing

In 3D and for NGLL=5 (Q4), for a regular hexahedral mesh there are:

125 GLL integration points in each element

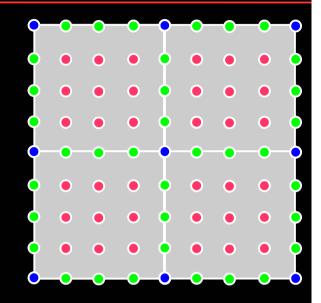
27 belong only to this element (21.6%)

54 belong to 2 elements (43.2%)

36 belong to 4 elements (28.8%)

8 belong to 8 elements (6.4%)

=> 78.4% of the GLL integration points belong to at least 2 elements



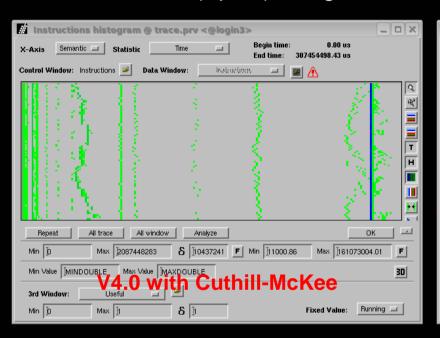
=> it is crucial to reuse these points by keeping them in the cache

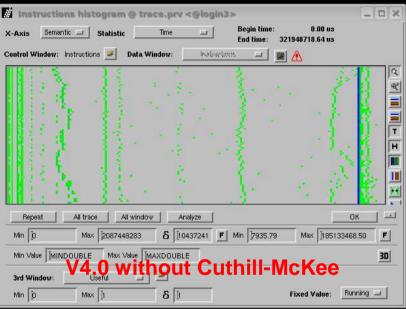
We use the classical reverse Cuthill-McKee (1969) algorithm, which consists in renumbering the vertices of the graph to reduce the bandwidth of the adjacency matrix

We gain a factor of 1.55 in CPU time on Intel Itanium and on AMD Opteron, and a factor of 3.3 on Marenostrum (the IBM PowerPC is very sensitive to cache misses)

#### Results for load balancing: instructions

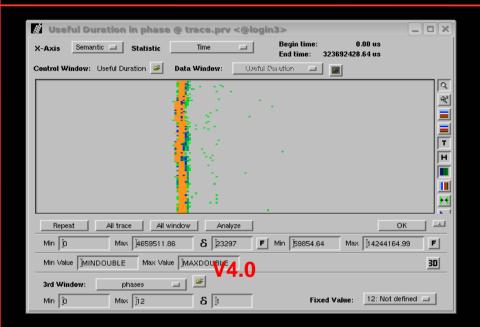
Analysis of parallel execution performed with Prof. Jesús Labarta in Barcelona (Spain) using his Paraver software package





- Number of instructions executed in each slice is well balanced
- ▶ Cuthill-McKee has almost no effect on that because we use high-order finite elements (of Q4 type), each of them fits in the L1 cache and for any such element we perform a very large number of operations using data that is already in L1

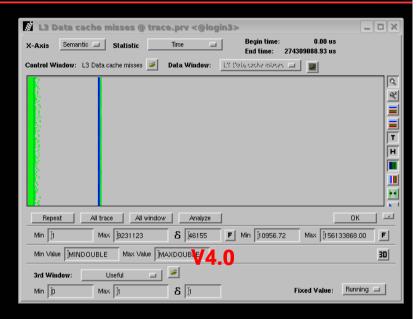
#### Results for load balancing: cache misses

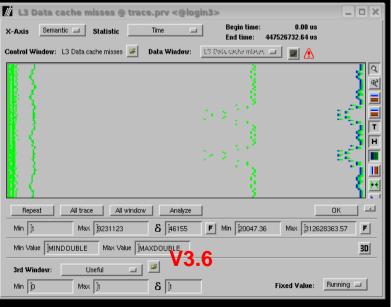


After adding Cuthill-McKee sorting, global addressing renumbering and loop reordering we get a perfectly straight line for cache misses, i.e. same behavior in all the slices and also almost perfect load balancing.

The total number of cache misses is also much lower than in v3.6

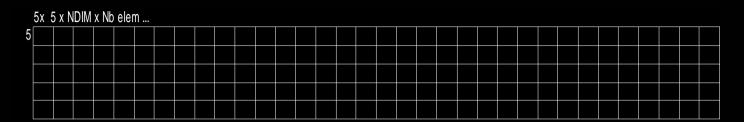
CPU time (in orange) is also almost perfectly aligned

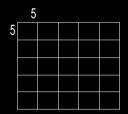




#### BLAS 3

(Basic Linear Algebra Subroutines)





Collaboration with Nicolas Le Goff (Univ of Pau, France)

Can we use highly optimized BLAS matrix matrix products (90% of computations)?

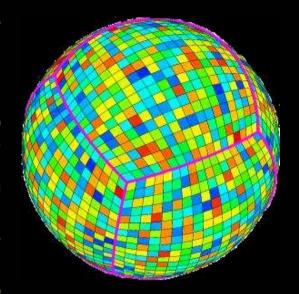
- For one element: matrices (5x25, 25x5, 5 x matrices of (5x5)), BLAS is not efficient: overhead is too expensive for matrices smaller than 20 to 30 square.
- ▶ If we build big matrices by appending several elements, we have to build 3 matrices, each having a main direction (x,y,z), which causes a lot of cache misses due to the global access because the elements are taken in different orders, thus destroying spatial locality.
- Since all arrays are static, the compiler already produces a very well optimized code.

=> No need to, and cannot easily use BLAS

=> Compiler already does an excellent job for small static loops

#### A very large run for PKP phases at 2 seconds

- ▼ The goal is to compute differential effects on PKP waves (collaboration with Sébastien Chevrot at OMP Toulouse, France)
- Very high resolution needed (2 to 3 seconds typically)
- Too big for current big machines (pangu: 4000 processors, Marenostrum 10000 processors) therefore we convert the mantle to acoustic instead of elastic and remove the crust because we are only interested in differential effects on P phases
- We keep an elastic anisotropic medium in the inner core only
- We can then design a mesh that is accurate down to periods of 2 seconds for P waves and that fits on 2166 processors (6 blocks of N x N slices, with N = 19)



- ► The mesh contains 126 billion points (the "equivalent" of a 5000 x 5000 x 5000 grid); We use 50000 time steps in 60 hours of CPU on 2166 processors on MareNostrum in Barcelona. Total memory used is 3.5 terabytes.
- Calculations with v4.0 (acoustic mantle, no crust) are finished, the code performed well and performance levels were very satisfactory; the analysis of the seismograms is under way

# Dec 26, 2004 Sumatra event



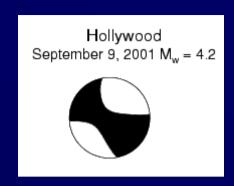


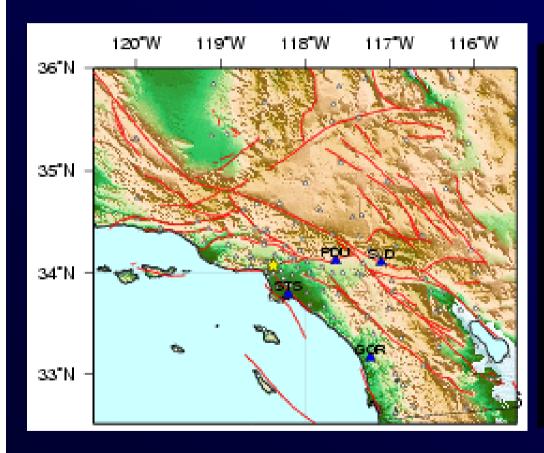
From Tromp et al., 2005

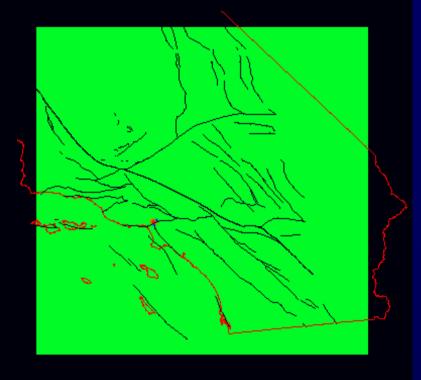
vertical component of velocity at periods of 10 s and longer on a regional scale

# Hollywood Earthquake

Small M 4.2 earthquake on Sept 9, 2001







Amplification in basin

# San Andreas – January 9, 1857

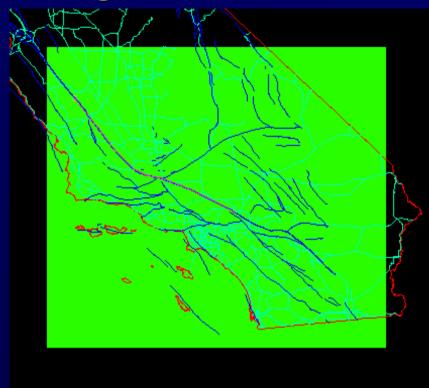


Vertical scale approximately 1 km

Carrizo Plain, San Andreas Fault, California, USA

# Earthquakes at the regional scale

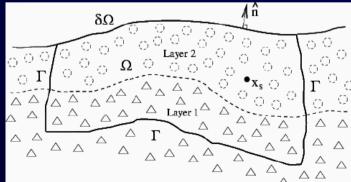


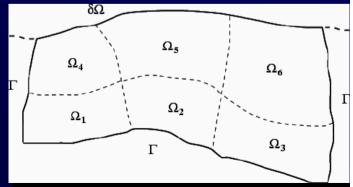


Carrizo Plain, USA, horizontal scale ≅ 200 m

I scale  $\cong 200 \text{ m}$ Scale approximately 500 km

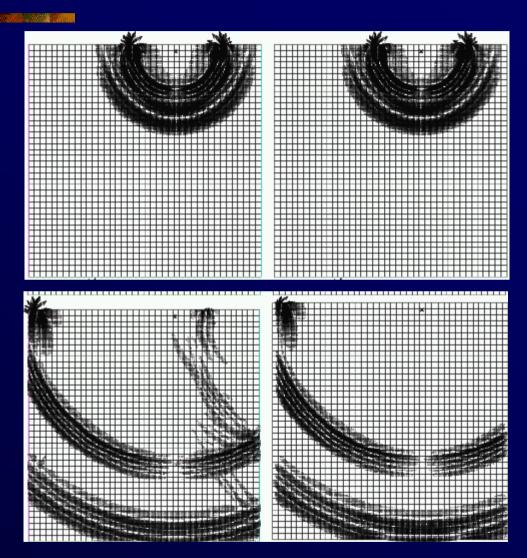
3D spectralelement method (SEM)





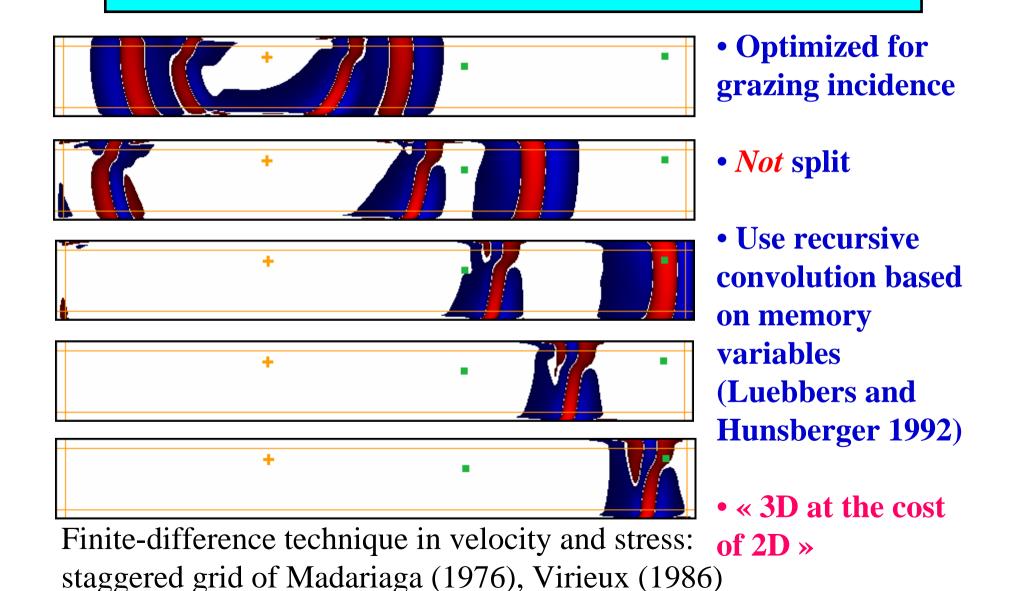
# Absorbing conditions

- Used to be a big problem
- Bérenger 1994
- INRIA (Collino, Cohen)
- Extended to second-order systems by Komatitsch and Tromp (2003)



PML (Perfectly Matched Layer)

#### Convolution-PML in 3D for seismic waves



## Future work

- ANR NUMASIS (2006-2009): optimize SPECFEM3D (among other codes) on NUMA machines (e.g. CEA Bull Tera10)
- Inverse problems (already done for the source by Liu et al 2004, but not for the model yet)
- Operto et al. for finite-differences  $\Rightarrow$  SEM in frequency? but full stiffness matrix  $\Rightarrow$  use MUMPS (ANR Solstice project)
- Altivec, CELL SuperScalar (Barcelona), GPGPU for the SPECFEM3D kernel (ENSEEIHT internship scheduled this year)
- Out-of-core for large-scale problems (talk by Jennifer Scott this morning, Abdou Guermouche in MUMPS etc)
- Grid computing (successful attempts on Egée): talk by Ronan Guivarch